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X-ray diffraction study of chalcopyrite CuFeS_2 , pentlandite $(\text{Fe, Ni})_9\text{S}_8$ and pyrrhotite Fe_{1-x}S obtained from Cu–Ni ore bodies

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Abstract. The x-ray diffraction (XRD) technique is applied to study five samples of Cu–Ni ore bodies, and it is shown that they contain chalcopyrite CuFeS_2 as the source of Cu, pentlandite $(\text{Fe, Ni})_9\text{S}_8$ as the source of Ni and pyrrhotite Fe_{1-x}S as a dominant compound. There are also other less dominant compounds such as bunsenite NiO , chalcocite Cu_2S , penrosite $(\text{Ni, Cu})\text{Se}_2$ and magnetite Fe_3O_4 . Using the XRD data obtained, we obtain the lattice parameters for tetragonal chalcopyrite as $a = 5.307(8) \text{ \AA}$ and $c = 10.38(2) \text{ \AA}$, that for cubic pentlandite as $a = 10.0487(13) \text{ \AA}$ and those for hexagonal pyrrhotite as $a = 6.882(4) \text{ \AA}$ and $c = 22.80(4) \text{ \AA}$. The lattice parameters relate to the compounds in their natural state.

1. Introduction

Cu and Ni are important metals technologically, and in this paper we apply x-ray diffraction (XRD) to study Cu–Ni ore bodies. Our XRD results show that the Cu–Ni ore bodies contain chalcopyrite CuFeS_2 as the source of Cu, pentlandite $(\text{Fe, Ni})_9\text{S}_8$ as the source of Ni and pyrrhotite Fe_{1-x}S as a dominant compound. There are also other less dominant compounds such as bunsenite NiO , chalcocite Cu_2S , penrosite $(\text{Ni, Cu})\text{Se}_2$ and magnetite Fe_3O_4 .

Apart from being a major ore of copper, chalcopyrite has several other interesting properties. The phonon and magnon dispersion of chalcopyrite has recently been reported by Harris *et al* [1] and its electronic structure has been studied by Kurmaev *et al* [2] and Fujisawa *et al* [3]. Chalcopyrite is a semiconductor with an optical absorption edge of about 0.5 eV [4], has a low electrical conductivity [5] and is an antiferromagnet with a very high Néel temperature of 823 K [6]. Also, semiconductors of the chalcopyrite structure have several applications, as photovoltaic devices, solar batteries, non-linear optical devices and luminescence diodes [7]. Compared to the case for chalcopyrite, there have been fewer studies done on pentlandite and pyrrhotite.

The plan of this paper is as follows. In section 2, we describe the experimental method which is based on XRD. Section 3 is devoted to results and discussion, and we show how the data are used to obtain the lattice parameters a , b and c for chalcopyrite, pentlandite and pyrrhotite. Concluding remarks are made in section 4.

|| Permanent address.

2. Experimental method

XRD was performed on five samples associated with Cu–Ni deposits from five different mines, operated by BCL Limited, Selebi Phikwe, in Botswana. Each sample was pulverized to obtain a powdered form. A computer-controlled Phillips Diffractometer, type PW3710, with an automatic divergence slit, and a Cu anode producing x-rays of wavelengths $\lambda_1 = 1.54060 \text{ \AA}$ and $\lambda_2 = 1.54439 \text{ \AA}$ were used, and data were collected in the mode where λ_2 was stripped off. The diffractometer was operating at 45 kV and 40 mA, and automatic routines allowed scanning for values of 2θ from 4° to 80° , using a step size of 0.02° and time per step of 0.02 s. Identification of each sample was achieved by comparison with a computerized database.

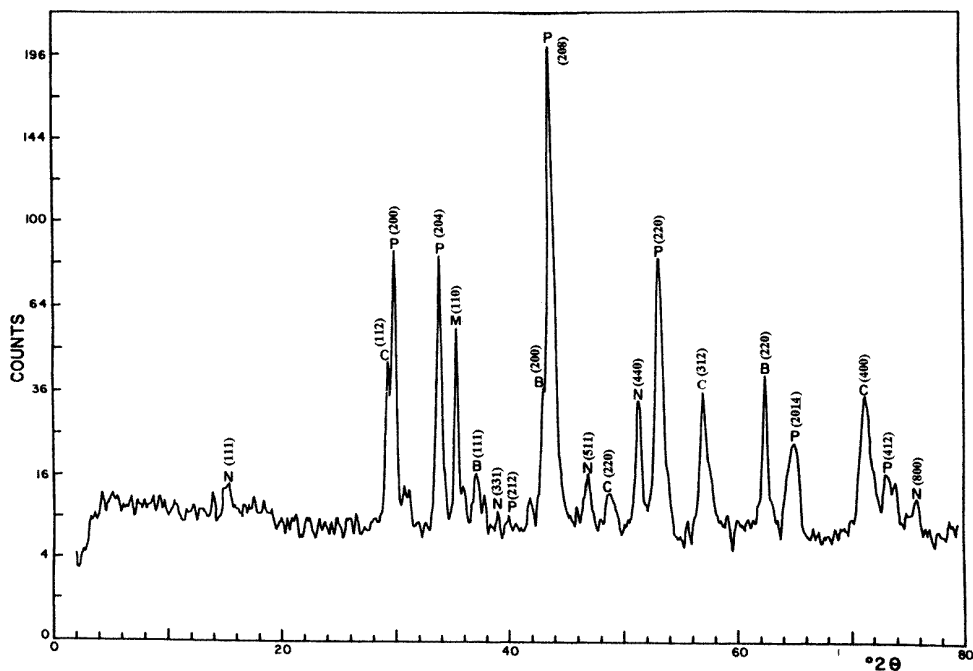


Figure 1. The XRD pattern of sample 1 of the Cu–Ni ore body, showing peaks for pyrrhotite (P), pentlandite (N), chalcopyrite (C), magnetite (M) and bunsenite (B).

3. Results and discussion

The experimental results for the five samples (to be referred to as 1 to 5) are presented as diffractograms in figures 1 to 5. By using Bragg's equation, the interatomic spacings, d , can be calculated from the angles which give diffraction peaks. The d -values are used to obtain the lattice spacings. All of the diffractograms show that the ore bodies contain chalcopyrite, pentlandite and pyrrhotite. Also, there are small amounts of magnetite and bunsenite in sample 1 (see figure 1), bunsenite and chalcocite in sample 2 (see figure 2), bunsenite, chalcocite and penrosite in sample 3 (see figure 3), magnetite and chalcocite in sample 4 (see figure 4) and bunsenite and chalcocite in sample 5 (see figure 5).

We shall, however, concentrate our discussion on the three major compounds present in the Cu–Ni ore bodies: chalcopyrite, pentlandite and pyrrhotite.

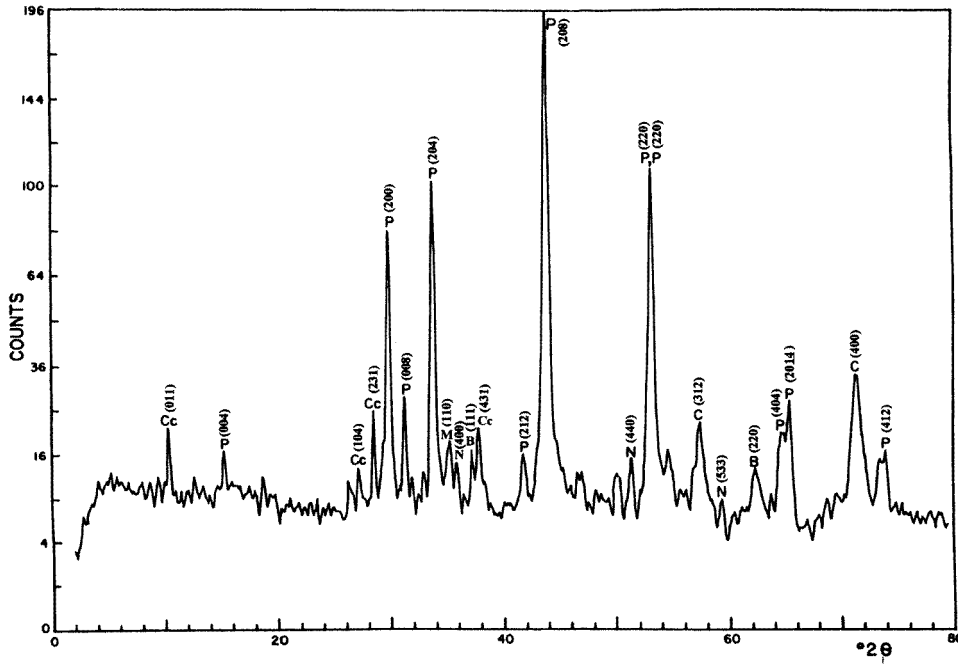


Figure 2. The XRD pattern of sample 2 of the Cu–Ni ore body, showing peaks for pyrrhotite (P), pentlandite (N), chalcopyrite (C), bunsenite (B) and chalcocite (Cc).

3.1. Chalcopyrite CuFeS_2

Chalcopyrite peaks are present in all of the diffractograms in figures 1 to 5. The highest diffracted intensity is from the (112) plane. In order to obtain the lattice parameters we proceed as follows. Chalcopyrite is a tetragonal crystal, and hence there are two unknowns, a ($=b$) and c , and therefore two equations are needed. Consider a set of values of d_1, h_1, k_1, l_1 , where d_1 is the interatomic spacing of atoms in the planes with Miller indices $(h_1 k_1 l_1)$, and d_2, h_2, k_2, l_2 , where d_2 is the interatomic spacing of atoms in the planes with Miller indices $(h_2 k_2 l_2)$. Then, for a tetragonal crystal, these values are related to the lattice parameters a and c by the following matrix equation:

$$\begin{bmatrix} (h_1^2 + k_1^2) & l_1^2 \\ (h_2^2 + k_2^2) & l_2^2 \end{bmatrix} \begin{bmatrix} 1/a^2 \\ 1/c^2 \end{bmatrix} = \begin{bmatrix} 1/d_1^2 \\ 1/d_2^2 \end{bmatrix}. \quad (1)$$

A solution to equation (1) obtained by using Cramer's rule gives the lattice parameters a and c for a tetragonal crystal as

$$a = \sqrt{\frac{(h_1^2 + k_1^2)l_2^2 - (h_2^2 + k_2^2)l_1^2}{(l_2^2/d_1^2 - l_1^2/d_2^2)}} \quad (2)$$

$$c = \sqrt{\frac{(h_1^2 + k_1^2)l_2^2 - (h_2^2 + k_2^2)l_1^2}{(h_1^2 + k_1^2)/d_2^2 - (h_2^2 + k_2^2)/d_1^2}}. \quad (3)$$

If either $l_1 = 0$ or $l_2 = 0$, we obtain

$$a = \begin{cases} d_1 \sqrt{(h_1^2 + k_1^2)} & \text{if } l_1 = 0 \\ d_2 \sqrt{(h_2^2 + k_2^2)} & \text{if } l_2 = 0. \end{cases} \quad (4)$$

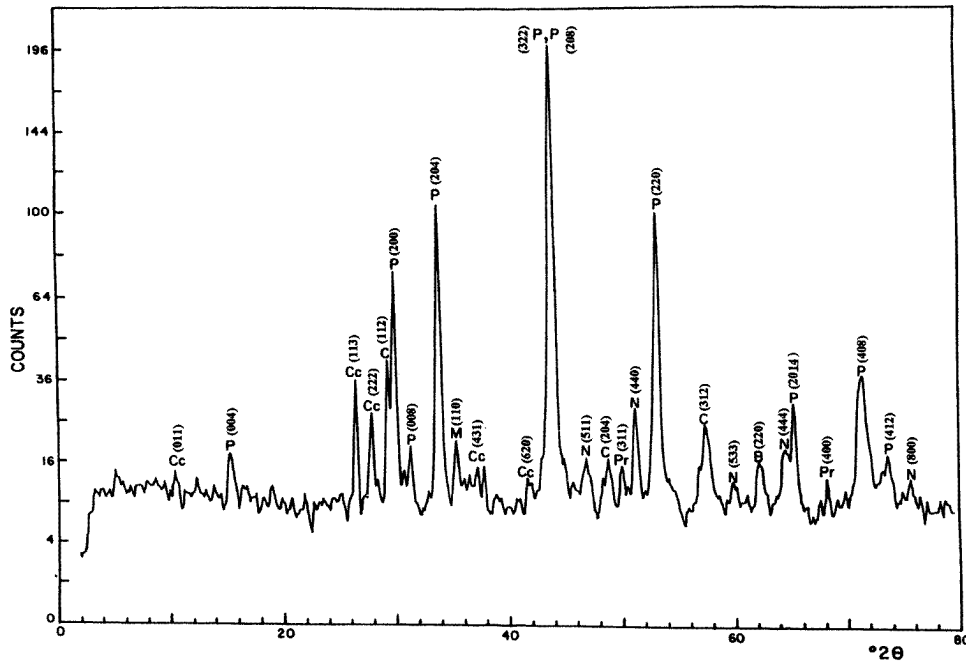


Figure 3. The XRD pattern of sample 3 of the Cu–Ni ore body, showing peaks for pyrrhotite (P), pentlandite (N), chalcopyrite (C), bunsenite (B), chalcocite (Cc) and penrosite (Pr).

If either $h_1 = k_1 = 0$ or $h_2 = k_2 = 0$, we obtain

$$c = \begin{cases} d_1 l_1 & \text{if } h_1 = k_1 = 0 \\ d_2 l_2 & \text{if } h_2 = k_2 = 0. \end{cases} \quad (5)$$

If $l_1 = l_2 = 0$ or $h_1 = k_1 = h_2 = k_2 = 0$, equations (2) and (3) should not be used. Using equations (2) and (3) and ten values from the XRD data, we obtain table 1, from which we find that the lattice parameters for chalcopyrite are $a = 5.307(8) \text{ \AA}$ and $c = 10.38(2) \text{ \AA}$. These values are comparable to those reported by Wyckoff ($a = 5.24 \text{ \AA}$, $c = 10.30 \text{ \AA}$) [8] and by Villars and Calvert ($a = 5.289 \text{ \AA}$, $c = 10.423 \text{ \AA}$) [9].

3.2. Pentlandite $(Fe, Ni)_9S_8$

Peaks due to pentlandite are illustrated in each of the figures 1 to 5. The highest diffracted intensity due to pentlandite occurs from the (440) plane. Pentlandite is a cubic crystal ($a = b = c$) and there is only one unknown, obtained as follows:

$$a = d\sqrt{h^2 + k^2 + l^2}. \quad (6)$$

Using equation (6) and ten values from the XRD data, we obtain table 2, from which we find that the lattice parameters for pentlandite are $a = 10.049(6) \text{ \AA}$. This value is comparable to that reported by Wyckoff ($a = 10.02 \text{ \AA}$) [8].

3.3. Pyrrhotite $Fe_{1-x}S$

Pyrrhotite peaks are distinct in all of the five diffractograms shown in figures 1 to 5, with the peak intensity in each of them occurring from the (208) plane. Pyrrhotite occurs in two

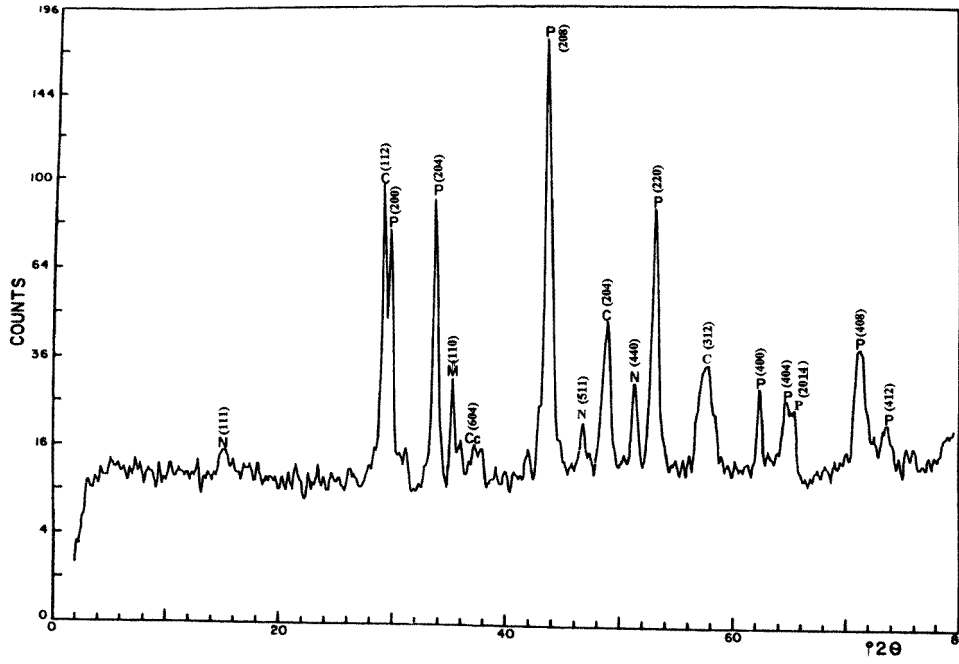


Figure 4. The XRD pattern of sample 4 of the Cu–Ni ore body, showing peaks for pyrrhotite (P), pentlandite (N), chalcopyrite (C), magnetite (M) and chalcocite (Cc).

Table 1. XRD values of d_1 , $(h_1k_1l_1)$, d_2 , $(h_2k_2l_2)$ used to calculate a and c for chalcopyrite.

Sample No	d_1 (Å)	$(h_1k_1l_1)$	d_2 (Å)	$(h_2k_2l_2)$	a (Å)	c (Å)
1	3.0366(2)	(112)	1.8656(1)	(220)	5.2767	10.4514
1	3.0366(2)	(112)	1.3239(0)	(400)	5.2956	10.3792
3	3.0457(2)	(112)	1.8570(1)	(204)	5.3221	10.3708
3	3.0457(2)	(112)	1.6024(1)	(312)	5.3295	10.3434
3	1.8570(1)	(204)	1.6024(1)	(312)	5.3287	10.3586
4	3.0498(2)	(112)	1.8556(1)	(204)	5.3524	10.3006
5	3.0442(2)	(112)	1.5925(1)	(312)	5.2851	10.4963
5	1.8720(1)	(220)	1.8556(1)	(204)	5.2948	10.4065
5	1.8720(1)	(220)	1.5925(1)	(312)	5.2948	10.3119
5	1.8556(1)	(204)	1.5925(1)	(312)	5.2892	10.4172
				Mean	5.307(8)	10.38(2)

symmetries: as Fe_{1-x}S in the hexagonal and monoclinic forms, and as Fe_7S_8 in the hexagonal form. Our XRD data show that the samples contain hexagonal Fe_{1-x}S . Using the same notation as in section 3.1, we obtain the following matrix equation for a hexagonal system ($a = b \neq c$):

$$\begin{bmatrix} \frac{4}{3}(h_1^2 + h_1k_1 + k_1^2) & l_1^2 \\ \frac{4}{3}(h_2^2 + h_2k_2 + k_2^2) & l_2^2 \end{bmatrix} \begin{bmatrix} 1/a^2 \\ 1/c^2 \end{bmatrix} = \begin{bmatrix} 1/d_1^2 \\ 1/d_2^2 \end{bmatrix}. \quad (7)$$

From equation (7), a and c for a hexagonal crystal are obtained as

$$a = \sqrt{\frac{4}{3} \frac{(h_1^2 + h_1k_1 + k_1^2)l_2^2 - (h_2^2 + h_2k_2 + k_2^2)l_1^2}{(l_2^2/d_1^2 - l_1^2/d_2^2)}} \quad (8)$$

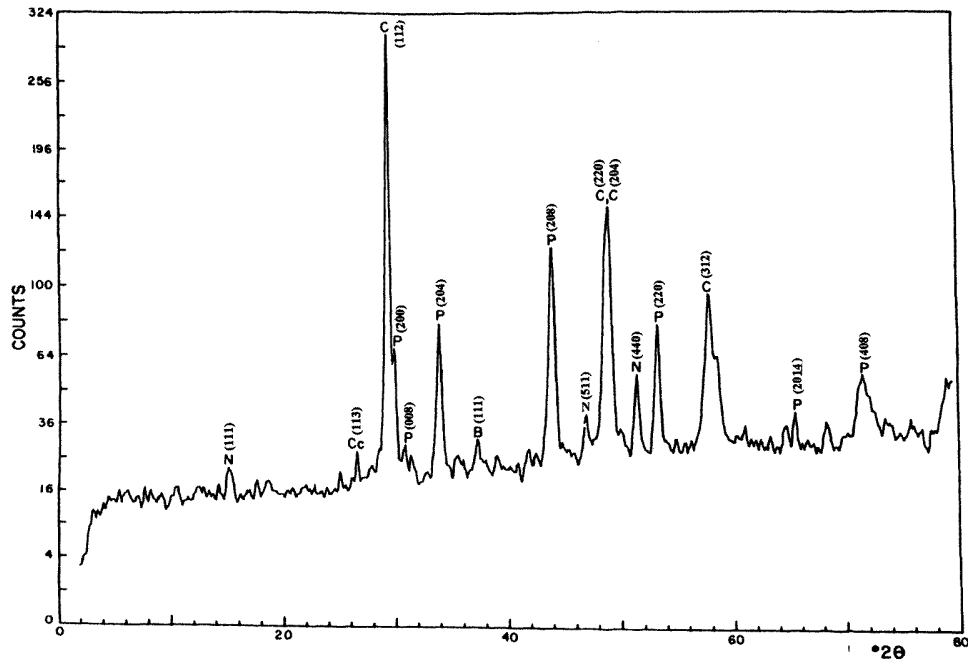


Figure 5. The XRD pattern of sample 5 of the Cu–Ni ore body, showing peaks for pyrrhotite (P), pentlandite (N), chalcopyrite (C), bunsenite (B) and chalcocite (Cc).

Table 2. XRD values of d and (hkl) used to calculate a for pentlandite.

Sample No	d (Å)	(hkl)	a (Å)
1	5.8110(11)	(111)	10.0649
1	2.3051(3)	(331)	10.0477
1	1.7763(2)	(440)	10.0483
2	1.7716(2)	(440)	10.0217
3	1.9291(2)	(511)	10.0239
3	1.7771(2)	(440)	10.0528
4	5.8015(11)	(111)	10.0485
4	1.9408(2)	(511)	10.0847
5	5.8034(11)	(111)	10.0518
5	1.7753(2)	(440)	10.0426
	Mean		10.049(6)

$$c = \sqrt{\frac{(h_1^2 + h_1k_1 + k_1^2)l_2^2 - (h_2^2 + h_2k_2 + k_2^2)l_1^2}{(h_1^2 + h_1k_1 + k_1^2)/d_2^2 - (h_2^2 + h_2k_2 + k_2^2)/d_1^2}} \quad (9)$$

If either $l_1 = 0$ or $l_2 = 0$, we obtain

$$a = \begin{cases} d_1 \sqrt{\frac{4}{3}(h_1^2 + h_1k_1 + k_1^2)} & \text{if } l_1 = 0 \\ d_2 \sqrt{\frac{4}{3}(h_2^2 + h_2k_2 + k_2^2)} & \text{if } l_2 = 0. \end{cases} \quad (10)$$

If either $h_1 = k_1 = 0$ or $h_2 = k_2 = 0$, c for the hexagonal system has the same form as that for the tetragonal system given in equation (5). If $l_1 = l_2 = 0$ or $h_1 = k_1 = h_2 = k_2 = 0$, equations (8) and (9) should not be used. Using equations (8) and (9) and ten values from the XRD data, we obtain table 3, from which we find that the lattice parameters for pyrrhotite are $a = 6.882(4) \text{ \AA}$ and $c = 22.80(4) \text{ \AA}$. These values are comparable to $a = 6.88 \text{ \AA}$ and $c = 22.90 \text{ \AA}$ reported in the *Diffraction File Data Book* [10].

Table 3. XRD values of d_1 , $(h_1k_1l_1)$, d_2 , $(h_2k_2l_2)$ used to calculate a and c for pyrrhotite.

Sample No	d_1 (\AA)	$(h_1k_1l_1)$	d_2 (\AA)	$(h_2k_2l_2)$	a (\AA)	c (\AA)
1	2.9714(2)	(200)	2.0641(1)	(208)	6.8622	22.9554
1	2.0641(1)	(208)	1.7236(1)	(220)	6.8944	22.8562
2	5.7251(8)	(004)	2.9835(2)	(200)	6.8901	22.9004
2	1.7209(1)	(220)	2.0619(1)	(208)	6.8836	22.8423
3	2.6373(2)	(204)	2.0547(1)	(208)	6.8779	22.7072
3	2.6373(2)	(204)	1.7184(1)	(220)	6.8736	22.7586
4	2.6517(2)	(204)	2.0717(1)	(208)	6.9020	22.9948
4	1.7208(1)	(220)	1.4411(0)	(404)	6.8832	22.6303
5	2.8427(2)	(008)	2.6396(2)	(204)	6.8827	22.7416
5	2.0510(1)	(208)	1.7177(1)	(220)	6.8708	22.6504
			Mean		6.882(4)	22.803(4)

4. Conclusions

We have used the XRD technique to study Cu–Ni ore bodies from five mines and have shown that the ore contains chalcopyrite as the source of Cu, pentlandite as the source of Ni and pyrrhotite as a dominant compound. There are also small amounts of bunsenite, chalcocite, penrosite and magnetite.

Using the XRD data, we have obtained the lattice parameters for tetragonal chalcopyrite as $a = 5.307(8) \text{ \AA}$ and $c = 10.38(2) \text{ \AA}$, for cubic pentlandite as $a = 10.049(6) \text{ \AA}$ and for hexagonal pyrrhotite as $a = 6.882(4) \text{ \AA}$ and $c = 22.80(4) \text{ \AA}$. The lattice parameters relate to the compounds in their natural state.

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